

# Quantifying large methane emissions from the Nord Stream pipeline gas leak of September 2022 using IASI satellite observations and inverse modelling

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## Abstract

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The sudden leaks from the Nord Stream gas pipelines, which began on 26<sup>th</sup> September 2022, released a substantial amount of methane (CH<sub>4</sub>) into the atmosphere. From the IASI instrument onboard EUMETSAT's MetOp-B, we document the first satellite-based retrievals of column-average CH<sub>4</sub> (XCH<sub>4</sub>) that clearly show the large CH<sub>4</sub> plume emitted from the pipelines. The data displays elevations greater than 200 parts per billion (ppb, ~11%) above observed background values ( $1882 \pm 21$  ppb). Based on the IASI data, together with an integrated

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mass enhancement technique and formal model-based inversions applied for the first time to thermal infrared satellite methane plume data, we quantify the total mass of CH<sub>4</sub> emitted to the atmosphere during the first two days of the leaks to be 219 - 427 Gg CH<sub>4</sub>. Substantial temporal heterogeneity is displayed in our model-derived flux rate, with three or four distinct peaks in emission rate over the first two days. Our range overlaps with other previous estimates, which were 75 – 230 Gg CH<sub>4</sub> and were mostly based on inversions that assimilated *in situ*

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observations from nearby tower sites. However, our derived values are generally larger than those previous results, with the differences likely due to the fact that our results are the first to use satellite-based observations of XCH<sub>4</sub> from the days following the leaks. We incorporate multiple satellite overpasses that monitored the CH<sub>4</sub> plume as it was transported across Scandinavia and the North Sea up to the evening of the 28<sup>th</sup> September 2022.

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We produced model simulations of the atmospheric transport of the plume using the Eulerian atmospheric transport model, TOMCAT, which show good representation of the plume location in the days following the leaks. The performance of simulated CH<sub>4</sub> mixing ratios at four nearby *in situ* measurement sites compared to the observed *in situ* values is mixed, which highlights the challenges inherent in representing short-term plume movement over a specific location using a model such as TOMCAT with a relatively coarse Eulerian grid. Our results confirm the leak of the Nord Stream pipes to clearly be the largest individual fossil fuel-related leak of CH<sub>4</sub> on record, greatly surpassing the previous largest leak (95 Gg CH<sub>4</sub>) at the Aliso Canyon gas facility in California in 2015-16.

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## 1. Introduction

Nord Stream is an offshore submerged pipeline network which carries natural gas from Russian facilities into Western Europe. The network is made up of two pipelines (NS1 and NS2), each originating in Russia and running through the Baltic Sea to Lubmin, Germany (Figure 1). NS1 has been operating since 2011 but the NS2 pipeline has not yet entered service, although it has carried natural gas. On 26<sup>th</sup> September 2022, multiple significant underwater gas leaks from these pipelines were detected by Nord Stream and the Danish Energy Agency, with apparently substantial gas emission through the water to the atmosphere (Danish Energy Agency, 2022). This was monitored by multiple national and international bodies over the following days. NS2 first began to leak on the morning of 26<sup>th</sup> September, from a location (15.41°E, 54.88°N) near the Danish island of Bornholm, whilst leaks were detected from NS1 at two more northerly locations (15.60°E, 55.54°N and 15.79°E, 55.56°N) later that day (leak locations marked with red stars in Figure 1). There were reports of explosions in the area around the times that these leaks were detected (e.g. GEUS, 2022) and the pressure in the pipelines underwent an abrupt and dramatic decrease, indicative of sudden ruptures in the pipes. Neither pipeline was transporting natural gas into Europe at the time, but both contained substantial quantities of gas, the vast majority of which is methane (CH<sub>4</sub>). This was released to the water and detected as large bubbles at the surface as it was further emitted into the atmosphere. Regions up to 0.7 km in diameter of rising gas bubbles were detected at the surface by *in situ* monitoring teams and by various satellite high-resolution imagers (e.g. Jia et al., 2022). The release of gas from the pipelines continued for a number of days before the Danish Energy Agency declared that the leaks had ceased on October 2<sup>nd</sup> 2022.

CH<sub>4</sub> is the second most significant greenhouse gas after carbon dioxide (CO<sub>2</sub>). Human-induced emissions of CH<sub>4</sub> have been responsible for 1.19 [0.81 – 1.58] Wm<sup>-2</sup> of anthropogenic effective radiative forcing since 1750 (net total of 2.72 [1.96 – 3.48] W m<sup>-2</sup>, Szopa et al. (2021)), with recent international agreements (UNFCCC, 2015; European Commission, 2021) having been put in place to urgently and significantly reduce CH<sub>4</sub> emissions for many countries. Recent satellite observations have shown that there are hundreds of CH<sub>4</sub> point source leaks worldwide contributing to direct anthropogenic emissions (e.g. Lauvaux et al., 2022). Growing levels of atmospheric CH<sub>4</sub> also adversely affect human health by contributing to increasing tropospheric ozone (West et al., 2006). A sudden large release of CH<sub>4</sub> into the atmosphere such as the one from Nord Stream could have significant consequences in terms of climate change and health. It is therefore important that the CH<sub>4</sub> emitted to the atmosphere during the Nord Stream leaks is accurately quantified. Various estimates, ranging from 75 to 230 Gg CH<sub>4</sub> (75,000 – 230,000 tonnes), have been suggested as to the quantity of CH<sub>4</sub> released to the atmosphere through assorted methodologies (see Jia et al. (2022); UNEP & IMEO (2023)).

Previous observational and modelling work (NILU, 2022; CAMS, 2022; NCEO, 2022; Jia et al., 2022) has shown that a plume of CH<sub>4</sub> originating from the location leaks was initially transported eastwards towards Finland's southern coast on 26<sup>th</sup> and 27<sup>th</sup> September, before a change in the wind direction then pushed it back out across Sweden and Norway and out into the North Sea to the north of Scotland late on the 27<sup>th</sup> and 28<sup>th</sup>. Significantly elevated near-surface CH<sub>4</sub> concentrations were briefly observed at a number of Integrated Carbon Observation System (ICOS) measurement towers in Scandinavia over the course of these three days, but there has been no direct satellite retrieval of downwind CH<sub>4</sub> concentrations available for the area to provide a more complete observation of the plume.

The Infrared Atmospheric Sounding Interferometer (IASI), on board EUMETSAT's MetOp-B satellite, is an across-track scanning thermal infrared sounder from which CH<sub>4</sub> distributions can be retrieved twice per day with high accuracy (Siddans et al., 2017). IASI's regular overpass times meant that it observed the area surrounding the CH<sub>4</sub> leak at approximately 09:30 and 21:30 local time each day. Thanks to favourable observing conditions, IASI observed enhanced CH<sub>4</sub> concentrations over the Baltic and the North Sea in the days following the detection of the Nord Stream leaks. We use this data, together with *in situ* observations from the ICOS network and an atmospheric chemical transport model, in order to quantify the total CH<sub>4</sub> emitted to the atmosphere from Nord Stream during the first two days of the leaks. This is the first time that plume flux inversions have been carried out using thermal infrared satellite data. Here we describe the results of this quantification and put into context the derived CH<sub>4</sub> contribution from these leaks compared both with previous similar large gas releases and with the global CH<sub>4</sub> budget.

Section 2 describes the IASI methane retrieval scheme used in this study, the CH<sub>4</sub> distributions retrieved from the satellite and the ICOS data. Section 3 describes the atmospheric model and the inverse modelling technique. We present our results in Section 4, before discussing their implications and concluding our discussion in Sections 5 and 6, respectively.

## 2. Observations

### 2.1 IASI retrievals

IASI is a cross-track-scanning Michelson interferometer (Blumstein et al., 2004) housed onboard the EUMETSAT polar-orbiting MetOp-B satellite, which was launched in 2012. Identical instruments are hosted on MetOp-A and -C, launched in 2006 and 2018, respectively, although MetOp-A is no longer operational. IASI provides daily global coverage with four circular footprints of approximately 12 km diameter at nadir, arranged in a 2 × 2 square grid of size 50 × 50 km. The IASI instrument measures upwelling thermal infrared radiation (TIR) with 8461 channels at 0.25 cm<sup>-1</sup> spectral resolution, ranging from 645 to 2760 cm<sup>-1</sup>. Observations are made at approximately 09:30 (descending node) and 21:30 (ascending node) local time each day. Column-average CH<sub>4</sub> distributions used here were retrieved using an updated version (v2.0) of a scheme developed originally for MetOp-A (Siddans et al., 2017), which has since been applied to MetOp-B (Knappett et al., 2022) and running in near-real time at the Rutherford Appleton Laboratory (RAL) in Oxfordshire, UK (<http://rsg.rl.ac.uk/vistool>). Updates included in the v2.0 scheme include improved representation of prior covariance, changes to spectroscopy in the radiative transport model, an updated elevation model and improvements to the representation of cloud, temperature and emissivity (Buchwitz et al., 2023). The v2.0 scheme retrieves CH<sub>4</sub> from measurements of its spectral signature in the 7.9 μm (1,260 cm<sup>-1</sup>) region (ν<sub>4</sub> fundamental vibration-rotation band). Vertical sensitivity generally peaks in the mid-upper troposphere since the spectral absorption signature is determined by temperature contrast with the surface. These data have previously been used for various studies of the atmosphere (e.g. Robson et al. (2020); Pope et al. (2021); Pimlott et al. (2022); Buchwitz et al. (2023)).

Elevated CH<sub>4</sub> mixing ratios were observed by IASI in the Baltic Sea above the leak sites on the morning of 26<sup>th</sup> September (Figure 2). However, cloudy conditions over much of Scandinavia and the North Sea meant that the plume was not detected during the evening overpass on 26<sup>th</sup> September, nor on the morning of 27<sup>th</sup> September. Large CH<sub>4</sub> mixing ratios off the northern coast of the UK on 27<sup>th</sup> September are likely unrelated to Nord Stream, although their source is unknown. Very high CH<sub>4</sub> concentrations were then detected over the North Sea off the west coast of Norway on the evening of September 27<sup>th</sup> and morning and evening of September 28<sup>th</sup>. On the morning of September 28<sup>th</sup>, in particular, a very distinct plume shape was detectable in IASI data, with areas of enhanced CH<sub>4</sub> around the northern and southern regions of the Norwegian coast. After that day, the plume became too diffuse to be distinguished from background concentrations. Retrieved column-averaged CH<sub>4</sub> (XCH<sub>4</sub>) enhancements within the plume on the morning of the 28<sup>th</sup> are up to 200 ppb (~11%), relative to the nearby background CH<sub>4</sub> mixing ratios of  $1882 \pm 21$  ppb (mean and standard deviation). The IASI retrievals documented here are the only satellite observations that captured a coherent XCH<sub>4</sub> plume from the Nord Stream leaks over the North Sea in the days immediately after the leaks began. On 30<sup>th</sup> September 2022, the GHGSat group's satellite constellation did capture a plume as it was emitted immediately above the leak location (GHGSat, 2022), although this was some days after the leaks began and by this point the emission rate was fairly small (~0.08 Gg hr<sup>-1</sup>). Although they operate at very high spatial resolution, GHGSat satellites retrieve only the CH<sub>4</sub> enhancement above the background, rather than total XCH<sub>4</sub>, and only targets specific sources. Meanwhile, Landsat-8-OLI and Sentinel-2B also detected enhanced CH<sub>4</sub> from high resolution images over the leak locations on 29<sup>th</sup> and 30<sup>th</sup> September (Jia et al., 2022), although these retrievals had large uncertainties associated with them.

## 2.2 ICOS network

Consistent *in situ* monitoring of CH<sub>4</sub> mixing ratios is carried out by the Integrated Carbon Observation System (ICOS) network (Levin et al., 2020; Heiskanen et al., 2022, <https://www.icos-cp.eu/>), a group of more than 140 monitoring sites located across Europe and Great Britain, including a number of measurement sites around southern Scandinavia. These sites measure greenhouse gas mixing ratios and fluxes in the atmosphere, ecosystems and oceans. The network includes 46 tall tower sites across 16 countries that measure greenhouse gas concentrations in the atmosphere, along with meteorological parameters. These include four sites near Scandinavia that continuously measure CO<sub>2</sub>, CH<sub>4</sub> and carbon monoxide (CO) mixing ratios at multiple heights between 10 m and 150 m above the surface. These are located at Birkenes, Norway (BIR, 8.3°E, 58.4°N, 219 metres above sea level (masl)); Hyltemossa, Sweden (HTM, 13.4°E, 56.1°N, 115 masl); Norunda, Sweden (NOR, 17.5°E, 60.1°N, 46 masl); and Utö, Finland (UTO, 21.4°E, 59.8°N, 8 masl). Sites are equipped with Picarro, Inc. G2401 cavity ring-down spectroscopy gas analysers, providing continuous CH<sub>4</sub> mixing ratios with a mean difference of  $0.2 \pm 0.8$  ppb compared to concurrent flask observations (Levin et al., 2020). The sites discussed here have inlets at heights between 10m and 150m above the ground (Hatakka et al., 2023; ICOS RI et al., 2023).

Significant enhancements of CH<sub>4</sub> (up to 770 ppb, or ~39%) were detected at each of these sites in the days following the Nord Stream leaks (Figure 3). We compare to the highest altitude inlet for each site, which ranges

between 57m and 150m above the ground across the four sites. UTO has only one inlet height. At the other locations, observed CH<sub>4</sub> mixing ratios can be quite different (up to 40 ppb at HTM and NOR and up to 300 ppb at BIR) across inlets and we choose the highest inlet height to attempt to reduce the impact of boundary layer mixing. There were relatively small CH<sub>4</sub> enhancements at UTO late on the 26<sup>th</sup> September, before larger enhancements were detected at NOR, HTM and finally BIR on the evening of the next day. The elevated concentrations at BIR were larger than at any other location. These observed values are consistent with the CH<sub>4</sub> plume from the leak being transported eastwards and then moving back westwards across Scandinavia before it was detected by IASI off the west coast of Norway on the 27<sup>th</sup> and 28<sup>th</sup> September. Here we used the data obtained at the ICOS locations for independent verification of our IASI-based analysis of the Nord Stream leaks.

### 3. Emission rate estimation methods and model description

We used two methods to estimate the total mass of CH<sub>4</sub> in the plume observed by IASI. We first applied an integrated mass enhancement (IME) technique, in tandem with Lagrangian model simulations, in order to estimate the total extra mass of CH<sub>4</sub> contained within the plume relative to local background concentrations. The Lagrangian model is used to inform the definition of the ‘plume’ and ‘background’ regions. This method has the advantage that, unlike formal inversions, it is not directly dependent on the accuracy of model transport to quantify the mass of CH<sub>4</sub> in the plume, but the main disadvantage is that it is not possible to exploit the averaging kernels (AKs) of the IASI retrievals to account for the vertical sensitivity of the derived XCH<sub>4</sub>, which peaks in the mid-upper troposphere. It also does not account for cloudy regions in which CH<sub>4</sub> is not retrieved. We therefore also employed a formal inverse modelling method based on simulation from a Eulerian chemical transport model which allowed us to model the plume directly and to take account of the satellite AKs.

The IME methodology used the Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model (Draxler and Hess, 1998) to produce a trajectory analysis which we combined with the XCH<sub>4</sub> data to determine boundaries for the enhanced CH<sub>4</sub> region due to the leaks. The HYSPLIT model was initiated with GFS meteorological data, with forward model trajectories starting at 1 km, 2km and 3km from 00:00 UTC on 26<sup>th</sup> September, running through to 00:00 UTC on 30<sup>th</sup> September. All three trajectories showed a similar pathway over the Baltic Sea, crossing Sweden during the morning of the 27<sup>th</sup> and reaching the Norwegian Sea by the 28<sup>th</sup> September. These trajectories, along with the IASI observations themselves, were used to define suitable enhanced XCH<sub>4</sub> regions and background regions, which represented the likely XCH<sub>4</sub> without the presence of the Nord Stream plume. The background regions were defined to the west of the calculated plume trajectories, at similar latitude ranges, away from the area affected by the leaks and over the ocean to preclude potential local sources of CH<sub>4</sub>. Background and enhancement regions are shown in Figure 2. The total additional CH<sub>4</sub> burden was calculated by computing the difference in the mean XCH<sub>4</sub> concentrations over the two regions and multiplying by the area. Estimates of the uncertainty were derived by perturbing the boundaries of the ‘background’ area chosen in each case with 4 scenarios, adjusting latitude- and longitude-box edges by  $\pm 1$  degree. We calculated estimates for the scenes observed on the morning of 26<sup>th</sup> September, the evening of the 27<sup>th</sup> and both the morning and evening of the 28<sup>th</sup>. The enhanced and background regions were allowed to vary

over time as the plume moved and dispersed across the North Sea. Multiple enhancement regions were permitted within a single overpass. Due to the cloud cover affecting our estimation of the ‘background’ XCH<sub>4</sub> in some cases, it is possible that these estimates include some sampling error that is difficult to quantify due to the cloud cover itself. We assume that this contribution to the uncertainty is small, however, since perturbing the boundaries of the background region does not affect large changes.

We also applied an atmospheric inversion technique to the IASI data to produce an optimised time-varying estimate of the emission rate for CH<sub>4</sub> from the leak. We used the global chemical transport model, TOMCAT (Chipperfield, 2006; Monks et al., 2017), to simulate the emission and transport of CH<sub>4</sub> from the location of the leak. TOMCAT has been used in a number of previous studies related to atmospheric CH<sub>4</sub> (e.g. McNorton et al., 2016, 2018; Wilson et al., 2016, 2021; Dowd et al., 2023), along with other atmospheric species. We ran the model at a horizontal resolution of  $1.125^\circ \times 1.125^\circ$ , which equates to approximately 65 km (east-west edges)  $\times$  125 km (north-south edges) at 60°N. There were 60 vertical levels from the surface up to 0.1 hPa. The model dynamical time step was 5 minutes. The model was forced by meteorological data from the European Centre for Medium-range Weather Forecasts (ECMWF) Operational analyses, regridded to the same horizontal and vertical resolution as the model grid. The meteorological data were read into the model every 6 hours, and linearly interpolated in time for each model time step. The initial conditions were produced from a previous forward simulation which ran up to 00:00 26<sup>th</sup> September 2022. Our simulation for the inversion ran from this time until 00:00 29<sup>th</sup> September 2022.

We simulated all non-plume-related CH<sub>4</sub> transport and chemistry as a separate tracer in the model, with all CH<sub>4</sub> fluxes from sources other than Nord Stream included in this background CH<sub>4</sub> tracer. Wetland emissions were taken from the WetCHARTs inventory (Bloom et al., 2017). Anthropogenic emissions were taken from the EDGAR v5 inventory (Crippa et al., 2020), whilst fire emissions were from GFED v4.1s (van der Werf et al., 2017). Emissions from all other sectors, the soil sink of CH<sub>4</sub> and the monthly mean offline atmospheric loss rates were as described in Wilson et al. (2021). Stratospheric loss rates due to O(<sup>1</sup>D) and chlorine are taken from a previous TOMCAT full chemistry simulation (Monks et al., 2017) and hydroxyl radical distributions are based on Spivakovsky et al. (2000). The enhanced XCH<sub>4</sub> observed by IASI is large compared to contributions from other sources, and the model run is short, so the effect of uncertainties from other sources and sinks of CH<sub>4</sub> should be minimal.

The emissions from the Nord Stream leak were treated as coming from point sources in the model (at 54.88°N, 15.41°E; 55.54°N, 15.60°E; and 55.56°N, 15.79°E), although these were instantly spread across the surface model grid cells containing the leaks. The southernmost leak was located near a model grid cell boundary in the longitudinal direction (at 15.2°E), so this leak was split equally between the two adjacent grid cells. This artificial instantaneous spreading out of the CH<sub>4</sub> from the leak will likely have some effect on the model’s representation of the plume movement but is unavoidable in a Eulerian model such as TOMCAT. Leak emissions during each 3-hour time window over the simulation were tagged as separate tracers to allow for independent scaling by the inversion (Figure S1). Figure 4 shows the TOMCAT column-averaged CH<sub>4</sub> at 08:30 UTC, the approximate IASI overpass time over the plume.

We tested three different *a priori* (prior) emission rate distributions. The first was a constant release rate of 4.17 Gg hr<sup>-1</sup> (4,170 tonnes hr<sup>-1</sup>) over the three days, emitting 300 Gg (300,000 tonnes) in total over this time. This value was chosen based on our initial test simulations and inversions (e.g. NCEO (2022)), and the results of previous studies (Jia et al., 2022). Additionally, this value proved to be a good compromise in producing simulated column mixing ratios approaching those seen IASI whilst not straying too far from the ICOS-observed values. The second distribution was an exponential decay with an e-folding lifetime of 24 hours, scaled to emit the same total CH<sub>4</sub> over the three days. The third distribution was taken from the results of a flow model described in Poursanidis et al. (2024). The flux rate derived in that work is based on the observed pressure change in the pipelines, the physical dimensions and depths of the pipes and various other factors. These prior emission rates are shown in Figure 5. We refer to these as the ‘constant prior’, the ‘decaying prior’ and the ‘modelled prior’ throughout this text. Although we scaled the constant and decaying priors to emit a total of 300 Gg CH<sub>4</sub> over the three days, the modelled prior was not scaled from the values provided in Poursanidis et al. (2024), emitting approximately 400 Gg CH<sub>4</sub> over the three days. The temporal variance of the modelled prior is closest to the most likely case, where the bulk of the emissions occurred after the NS1 leaks began on the evening of the 26<sup>th</sup> September.

We carried out Bayesian inversions based on analytical calculation of an *a posteriori* (posterior) leak emission rate based on finding the minimum of a cost function as in Tarantola and Valette (1982). We optimised the mean flux from the leak locations for each 3-hour window throughout the simulation and the mean background XCH<sub>4</sub>, giving 25 optimised values in total. The mean background XCH<sub>4</sub> was given a prior uncertainty of 1%, equal to approximately 18 ppb, and was changed very little by the inversion. All other sources and sinks were kept unchanged. We assimilated only the data from the morning of September 28<sup>th</sup> (Figure 2e), since this overpass detected the most coherent and extensive observation of the plume. We either assimilated all observations made that morning (3980 individual retrievals, denoted ‘all’), or retrievals only within the region bounded by the longitudes 3.5°W and 9.8°E and the latitudes 58.7°N and 70.0°N, the region that contained the main mass of the plume on the morning of 28th September (905 individual retrievals, denoted ‘plume’, see Figure 6a for region definition). The AK associated with each IASI sounding was applied to the corresponding TOMCAT methane profile. Due to the small number of variables that we optimise, and the relatively small number of observations included, the posterior solution can be solved for directly, as has been done previously using TOMCAT (e.g. McNorton et al., 2018; Claxton et al., 2020). See Supplementary Material and those references for more detail of the inversion method.

We tested both the assumption that the Gaussian emissions uncertainties during each 3-hour window were uncorrelated with each other (nocorr), and that consecutive emission windows had uncertainties with correlations of 0.7 (corr). This value was chosen in order to impose a fairly strong correlation between emission windows but proved to have little impact on results during emission windows that were well-constrained by observations (See Figure 5). We tested prior uncertainties of both 100% and 50% (denoted 1.0 $\sigma$  and 0.5 $\sigma$ ). Additionally, instead of optimising against the full set of individual IASI retrievals, we tried optimising only the single mean XCH<sub>4</sub> value within the bounded region described above (denoted ‘regional mean’) and finally we

optimised against mean XCH<sub>4</sub> values at  $3^\circ \times 3^\circ$  horizontal resolution (see Supplementary Material). These simulations were intended to account for discrepancies between the simulated location of the plume compared to the observed location due to transport errors. In total we therefore carried out 48 different inversions based on different prior emission distributions, sets of assimilated data and assumptions regarding prior uncertainties (see Table S2 for details of the inversions). In all inversions, the uncertainty on the retrievals was set at 30 ppb and were assumed to be uncorrelated with each other. This value is more conservative than the estimated individual IASI sounding uncertainty ( $\sim 20$  ppb), in order to attempt to account for uncertainties from the model transport. We applied the IASI averaging kernels to represent the satellite's vertical sensitivity in the simulated column average values. The matrices were inverted using LU decomposition methods.

For comparison of our results with the ICOS CH<sub>4</sub> observations, we interpolate the simulated prior or posterior mixing ratios from all tracers to the corresponding latitude, longitude and inlet heights of the ICOS sites, before adding them together to produce simulated time series of CH<sub>4</sub> at each of the four sites. At each site, we compared to the observational data obtained at the highest inlet height available, to attempt to reduce the influence of boundary layer mixing.

## 4. Results

### 4.1 Integrated Mass Enhancement (IME) results

The IME method yielded various total mass estimates for each of the overpass times during the first three days of the leak. The results are shown in Table 1. The first estimate of  $30 \pm 1$  Gg CH<sub>4</sub> is from an overpass that occurred only a few hours after the first leak began. Assuming that the leak commenced at 02:00 local time and that IASI was able to view most of the leaked CH<sub>4</sub> during this overpass, this implies a mean emission rate of  $\sim 4$  Gg hr<sup>-1</sup> during that time. However, many nearby areas were obscured by cloud, so it is likely that IASI could not view all of the CH<sub>4</sub> emitted during these initial hours. The estimate at this time is therefore likely to be an underestimate of the total CH<sub>4</sub> release.

No plume was visible for the next 36 hours, before what was quite likely only a partial view of the plume obtained on the evening of 27<sup>th</sup> September on the west coast of Norway. The total CH<sub>4</sub> mass within the small observed section of this plume was  $16 \pm 1$  Gg, the low value likely due to much of the plume being obscured by cloud. A very clear view of the plume, which by this point was beginning to split into northern and southern sections, on the morning of 28<sup>th</sup> September yielded an inferred total of  $161 \pm 4$  Gg of CH<sub>4</sub>. Finally, a total enhancement of  $77 \pm 2$  Gg was calculated for the evening of the 28<sup>th</sup>.

Analysis of these values is complex for two reasons. First, the effect of the IASI instrument's vertical sensitivity through application of AKs has not been taken into account. The consequences of this are hard to quantify as they depend on the vertical sensitivities of IASI both within the plume and in the background regions, and the



actual vertical distribution of the CH<sub>4</sub> within the column in those regions. Using the TOMCAT model to compare the total column values in the plume with and without the AKs applied indicates that the error due to this effect may be up to 4% of the total column value and 37% of the mass enhancement value, although this relies on the accuracy of the model's vertical transport. Second, it is possible, and on some overpasses likely, that not all of the CH<sub>4</sub> emitted from the leak was viewed by the satellite, which would introduce a negative bias to the results. It is clear that for the majority of the IASI overpasses, at least some part of the plume is unfortunately obscured by cloud. The only clear view of the plume is obtained on the morning of the 28<sup>th</sup> September, as confirmed by the TOMCAT simulations. We therefore suggest that the IME-related estimates from the other overpasses are likely to be underestimates of the total CH<sub>4</sub> released by Nord Stream, and we base any conclusions on the estimate obtained on the morning of September 28<sup>th</sup>.

**Table 1: Integrated mass enhancement (Gg CH<sub>4</sub>) calculated from the Nord Stream plume observed by IASI over three days in September 2022. Also included are the defined enhancement region and background region boundaries. Overpass times with 'N/A' stated are for overpasses when the satellite's view of the CH<sub>4</sub> plume was totally obscured by cloud.**

Approximate local overpass time (hh:mm DD/MM/YY)	Enhancement region boundaries	Background region boundaries	Total derived CH <sub>4</sub> mass enhancement (Gg)
09:30 26/09/22	53°N – 56°N; 13°E – 17°E	64°N – 70°N; -4°E – 0°E	30 ± 1
21:30 26/09/22	N/A	N/A	N/A
09:30 27/09/22	N/A	N/A	N/A
21:30 27/09/22	64°N – 66°N; 8°E – 10°E	64°N – 70°N; -4°E – 0°E	16 ± 1
09:30 28/09/22	1) 59°N – 63°N; -2°E – 4.5°E 2) 63°N – 70°N; 4°E – 7°E 3) 66°N – 71°N; -12°E – -8°E	64°N – 70°N; -12°E – -8°E	161 ± 4
21:30 28/09/22	1) 68°N – 72°N; -8°E – 4.5°E 2) 59°N – 63°N; 1°E – 4°E	64°N – 68°N; -12°E – -8°E	77 ± 2

## 4.2 Inversion results

Figure 4 shows the development of the simulated Nord Stream plume in the TOMCAT model over the first three days of the leak, assuming constant emission rates during this time. The plume initially moves northwards and eastwards during the first day. Over the following two days the plume is transported rapidly westwards across Sweden and Norway, before emerging over the North Sea at a similar time and location as indicated by the satellite observations. The plume becomes quite diffuse by the evening of 28<sup>th</sup> September.

The prior emissions, in both the ‘constant’ and ‘decaying’ configurations, underestimate the observed XCH<sub>4</sub> in the plume region on the morning of 28<sup>th</sup> September (Figure 6 and Figures S2 – S4). The simulated location of the northern section of the plume is also slightly east of the observed location. Using the ‘modelled’ prior, high XCH<sub>4</sub> values like those observed by IASI are produced, but in an inaccurate location, to the south and west of the observed plume (Figures S5 and S6). These discrepancies are likely due to a combination of underestimation of the initial leak rate, errors in the timing of the peak emissions in the prior and model transport errors. It is possible that the meteorological analyses used in the model and the vertical mixing parameterisation in TOMCAT combine to produce small errors in the simulated plume position. Figure 7 shows the total posterior emissions over the first two days of the leaks. In all cases, the posterior emissions are larger than 200 Gg produced by the ‘constant’ and ‘decaying’ prior emissions. We report totals for only the first two days, as the observations provided by IASI on the morning of 28<sup>th</sup> September do not constrain emissions on the third day. The mean posterior emission total for these two days is  $299 \pm 50$  Gg (here the reported uncertainty represents the standard deviation across the mean posterior values). The mean posterior total is  $280 \pm 35$  Gg when omitting the ‘regional mean’ inversions where only the mean CH<sub>4</sub> value is optimised. These values are close to the value used in the ‘modelled’ prior, based on the work by Poursanidis et al. (2024). However, there is significant variation in the individual posterior totals, which range between  $219 \pm 23$  Gg and  $427 \pm 69$  Gg, depending on the assumptions made (here the uncertainty represents the derived posterior uncertainty from the individual inversion). Total posterior emissions are consistently smaller when applying the ‘decaying’ prior than with the ‘constant’ or ‘modelled’ priors. Using the ‘modelled’ prior, emissions larger than 400 Gg are derived when only the regional mean is optimised. When optimising using the  $3^\circ \times 3^\circ$  average XCH<sub>4</sub>, posterior results are consistent with the other methods, ranging from 211 – 294 Gg (see Table S1).

When the inversion optimises the model using the individual IASI retrievals for the ‘constant’ and ‘decaying’ priors, the position of the northern section of the plume is improved (moved further west), similar to the observations (Figure 6 and Figures S2 – S4), and simulated XCH<sub>4</sub> is increased. However, the XCH<sub>4</sub> still remains lower than the observed values. When the regional mean is optimised, the magnitudes of the simulated XCH<sub>4</sub> values are much improved, but the position of the largest values is not improved relative to the IASI observations. Using the ‘modelled’ prior, optimising against the individual retrievals increases the XCH<sub>4</sub> in the northern section of the plume. However, when optimising against the regional mean, the posterior plume location remains too far south and west. The remaining errors in the model representation of the plume are likely due to: i) errors in the ECMWF meteorological data, which might be improved through use of reanalyses rather than the operational analyses; ii) biases in the model transport parameterisations, particularly for vertical

mixing, leading to incorrect simulated vertical distribution of the plume; and iii) uncertainties produced due to the instantaneous mixing of the leak emissions across model grid boxes.

The three-hourly posterior emission rates display significant variation over the first two days of the leaks (Figure 5). When the ‘constant’ or ‘decaying’ prior are used, there are three peaks in the posterior flux rates – the first during the early afternoon on the 26<sup>th</sup> September, and two more smaller peaks during the morning and afternoon of the 27<sup>th</sup>. Using the ‘modelled’ prior produces four peaks. There are low emission rates between these times. This temporal variation is consistent across all inversions, including, to some extent, when only the regional mean XCH<sub>4</sub> is optimised (Figure S5). The posterior emissions are far outside of the prior uncertainty during peak flux rates and, in fact, are below zero during the night of 26<sup>th</sup>. This negative flux is also suggestive of model transport errors or biases in the background concentration. Unless temporal error correlations are included for the prior flux in an inversion, emissions during the third day are not constrained.

Figure 3 includes the CH<sub>4</sub> mixing ratios observed at the four ICOS sites for 26<sup>th</sup> – 29<sup>th</sup> September, and the prior and posterior model values at those locations. The largest observed CH<sub>4</sub> enhancements above the background concentrations were at BIR (~770 ppb), with enhancements of ~340 ppb at NOR and HTM and much smaller enhancements of less than 60 ppb at UTO. The prior model simulations are close to the observations at UTO. At BIR, the peaks in the prior simulations occur around 3 hours too early. The constant and decaying priors underestimate the magnitude of the peak CH<sub>4</sub> whilst the modelled prior overestimates the magnitude. The timing of the peak in the prior simulation at NOR is similarly early and the magnitude is 200 – 700 ppb too high for the constant and decaying priors, but well-captured by the modelled prior. Finally, the model performance at HTM is poor, with very large simulated values, likely due to the site’s location relative to the model grid boundaries and the fast spreading of the leak emissions both leading to excessive influence from CH<sub>4</sub> directly from the leaks. The performance of the IASI-based posterior emissions at the ICOS sites is mixed. Peak CH<sub>4</sub> at each site has generally remained the same or increased. Posterior values at HTM have significantly increased, whilst performance at UTO has changed little. The posterior performance is improved at BIR, both in terms of the timing and magnitude of the peak, but at NOR, the posterior peaks remain much too large.

## 5. Discussion

The range of estimates from both of the methodologies that we applied to estimate the total CH<sub>4</sub> emitted from the Nord Stream leaks using IASI retrievals of XCH<sub>4</sub> produced values greater than 200 Gg, with some estimates reaching more than twice that value. A leak of this magnitude is by far the largest individual anthropogenic leak of CH<sub>4</sub> to the atmosphere on record, at least twice as large as the previous largest emission event in Aliso Canyon, California in 2015-2016 (97 Gg, Conley et al. (2016)). That leak was from a ruptured injection well pipe at a gas storage facility near Los Angeles and continued for more than three months.

The magnitude of the Nord Stream leaks is highly significant on a global scale – when considered over a short period. Total global CH<sub>4</sub> emissions from fossil fuels amounted to 108 Tg in the year 2017 (Saunio et al. (2020), top-down estimate), approximately 300 Gg day<sup>-1</sup>. Our mean estimate from the Nord Stream leaks over two days

is therefore approximately equivalent to an extra day's emissions from global fossil fuel sources (although it should be noted that daily emissions are likely larger today than they were in 2017). However, in the context of annual anthropogenic CH<sub>4</sub> emissions (~364 Tg yr<sup>-1</sup>), the Nord Stream leaks contributed only an extra 0.08%, and increased the annual global total CH<sub>4</sub> emissions from all sources (~600 Tg yr<sup>-1</sup>) by just 0.05%. Chen and Zhou (2023) calculated that a leak from Nord Stream of magnitude 220 Gg would have a negligible warming effect on the climate ( $1.8 \times 10^{-5} \text{°C}$  over a 20-year period) and our larger emission estimates would have a correspondingly small effect.

IASI had its best view of the plume during the morning of 28<sup>th</sup> September 2022, and we base our best estimate of the total CH<sub>4</sub> leaked to the atmosphere during the preceding two days on the observations made at that time. Our IME method produced a value of  $161 \pm 4$  Gg CH<sub>4</sub> from those retrievals, whilst our TOMCAT inversion results produced a range of 219 - 427 Gg, with a mean of  $280 \pm 35$  Gg when optimising the model based on comparisons to individual retrievals. The consistency between the results produced using the two methods is therefore poor, with the IME value approximately 40% smaller than the inversion mean. This is likely due in part to the fact that the IME method does not take account of IASI's vertical sensitivity with results being affected by up to 37% by this. The effect of missing IASI data due to cloud cover on the estimated IME value (and to a lesser extent, on the inversions) is also difficult to quantify. It should be noted that the posterior XCH<sub>4</sub> produced by TOMCAT is in most cases, still underestimating compared to the IASI observations, indicating that some posterior estimates may be underestimates.

We investigated the vertical structure of the simulated plume, together with the vertical sensitivity of XCH<sub>4</sub> retrievals based on the IASI AKs (Figure S6). This shows that the northern and southern sections of the plume during the morning of 28<sup>th</sup> September (defined as 66°N – 71°N, -5°E – 6°E and 59°N – 63°N, 0°E – 7°E, respectively) have different vertical structures in the model. The northern section has high near-surface CH<sub>4</sub> mixing ratios from the leaks, which remain relatively constant as with altitude before decreasing until there is no influence from Nord Stream above 500 hPa (~5.5 km). In this case, the majority of the leak-related CH<sub>4</sub> is located beneath the peak IASI vertical sensitivity indicated by the AKs. Meanwhile, in the southern section, the CH<sub>4</sub> contribution from the leak is smaller, but peaks higher up, at approximately 600 hPa (~4 km), around the same region as the peak satellite sensitivity. If the vertical distributions produced in the model are correct, this indicates that the observed XCH<sub>4</sub> in the northern and southern sections of the plume, whilst displaying similar XCH<sub>4</sub> values, are in fact due to very different relative CH<sub>4</sub> contributions within the column. If the simulated vertical distributions are correct, it is likely that the IME method underestimates the CH<sub>4</sub> mass in the northern section of the plume whilst overestimating it in the southern section.

The interpretation that the inversion-derived values are low is complicated by performance of the posterior simulations at the ICOS site locations (Figure 3). The high values observed at BIR are captured well in the posterior, but the model overestimates overserved values significantly at NOR. This is also true when using the prior emissions, however, indicating that model performance might not be accurate at NOR. In the model, the HTM site is located in a grid box next to the one into which the Nord Stream CH<sub>4</sub> is emitted, and the comparison there is likely negatively and unrealistically affected by this. In fact, an inversion based only on

assimilating the ICOS observations, without the IASI data, produces a much smaller posterior total emission ( $88 \pm 13$  Gg, Figure S7). We hypothesise that our Eulerian model's representation uncertainty is large when  
465 simulating the movement of a large distinct plume over fixed point measurement locations, especially at the  
resolution used here. In addition, the model's representation of the detailed vertical structure of the plume is key  
for such comparisons. The use of a high-resolution regional model, a nested grid, or a Lagrangian model might  
produce better comparisons at the ICOS sites.

Our IASI-based estimates are consistently larger than estimates produced by others using different observational  
470 datasets. Previous estimates issued by our team and by other groups were produced quickly in the weeks  
immediately following the leaks, and we have here attempted to probe the sensitivity of our results to chosen  
methodologies and assumptions about the leaks and observational data. Based on ICOS observations, satellite-  
based imaging spectrometer data and multiple Lagrangian models, Jia et al. (2022) calculated a total flux of  $220 \pm 30$  Gg  $\text{CH}_4$  over three days of leaks, which itself was larger than many estimates published by various groups  
475 using a range of methods and datasets (CAMS, 2022; NILU, 2022; UNEP & IMEO, 2023). The temporal  
variation of emissions produced by Jia et al. (2022) showed some similarity to our own results, with the peak  
emission rate occurring during the night of 26<sup>th</sup> -27<sup>th</sup> September, more than 24 hours after the first of the leaks  
began, but only a few hours after three of the four leaks started. They also computed the mass of  $\text{CH}_4$  that was  
released from the pipelines based on pipeline dimensions and the change in gas pressure within the pipes,  
480 calculating a value of 230 Gg. This value, along with their calculated emission value, is smaller than the  
majority of our emission estimates, although a subset of our results is consistent with their value. It remains  
important to investigate the roots of the apparent discrepancies between our IASI-derived estimates and those  
produced via other means.

485 The resolution used by TOMCAT in this case (approximately  $1^\circ \times 1^\circ$ ), is fairly coarse for capturing the  
movement of the plume over the ICOS sites, and results will be affected by the artificial instantaneous spreading  
of the point source emissions over the comparatively large model grid cells. The coarse resolution likely has less  
impact on the model-satellite comparisons overall, however. We can employ Eulerian models with higher  
resolution, and/or Lagrangian plume models, to attempt to better represent the plume's distribution in  
490 comparison with IASI. The effect of the meteorological data used in the models can also be assessed through the  
use of reanalyses from ECMWF or other meteorological datasets. The operational meteorological analyses used  
here are updated by ECMWF during reanalysis through assimilation of satellite and *in situ* observations, which  
might result in better consistency between the simulated and observed plume. The uncertainty induced by the  
emission inventories should be small compared to the observed plume-related concentrations during a short  
495 simulation such as this one, but the initial conditions could introduce biases between the model and satellite. We  
attempt to account for this through inclusion of the background  $\text{XCH}_4$  in the inversion state vector, but further  
investigation into the effect of the initial conditions is warranted. In addition, investigation into the model's  
representation of plume uplift above the  $\text{CH}_4$  release to the atmosphere might be a key uncertainty, since it  
determines layer height and therefore the horizontal wind field to which the simulated plume is exposed.

## 500 6. Summary and Conclusions

We have produced the first clear satellite retrievals of column average methane that capture the CH<sub>4</sub> emitted into the atmosphere from the Nord Stream gas leaks in late September 2022. The IASI instrument, onboard the satellite MetOp-B, produced retrievals displaying strongly enhanced XCH<sub>4</sub> at the leak locations on the morning of 26<sup>th</sup> September, before large widespread enhancements were seen over the North Sea during 28<sup>th</sup> September. The satellite data retrieved for that day allowed us to employ two methods to quantify the CH<sub>4</sub> leaked to the atmosphere from the Nord Stream leaks during the first two days.

Our integrated mass enhancement calculations produced total emissions of  $161 \pm 4$  Gg CH<sub>4</sub>, although this method cannot take account of the satellite instrument's vertical sensitivity, which peaks in the mid-upper troposphere, and cannot account for regions of enhanced CH<sub>4</sub> that are not observed due to clouds. We also used formal Bayesian inversion methods, using the TOMCAT atmospheric chemical transport model, to quantify the emissions based on the observations made on the morning of 28<sup>th</sup> September. This is the first time that plume flux inversions have been carried out using thermal infrared satellite data. Here, we investigated the effect of a range of assumptions within the inversion, including the prior distribution of the emissions, the related prior uncertainties and the way that observations are assimilated. We calculated total emissions between 219 and 427 Gg. The mean over all inversions is approximately  $299 \pm 50$  Gg, whilst the mean over the inversions that optimise against individual IASI retrievals is  $280 \pm 35$  Gg. All of our results imply that the Nord Stream leaks were by far the largest recorded individual anthropogenic leak of CH<sub>4</sub> to the atmosphere.

Our estimates are larger than previous values given for the Nord Stream leaks, produced using alternative observational data. There is variable performance when comparing our posterior results to *in situ* observations made in the region, and more work is necessary to discern to what extent this is due to errors in the flux estimates produced from the satellite data and how much is due to poor model plume representation at some tall tower locations. Our ability to monitor, simulate and quantify leaks of GHGs and pollution events such as this one is continuously improving, aiding our ability to mitigate the human influence on the atmosphere. It is also clear from this study that thermal infrared instruments such as IASI, which have peak sensitivity high in the troposphere, are able to provide more information concerning surface events such as the Nord Stream leaks than might have been appreciated previously. In any case, whilst this particular event remains highly significant locally over a short time period, the effect of these emissions, by themselves, is very small in terms of both the global atmospheric CH<sub>4</sub> budget and the climate.

#### Data Availability

MetOp-B IASI methane observations up to March 2021 are available on the Centre for Environmental Data Analysis (CEDA) long-term data archive (Knappett et al., 2022). More recent data, including the near-real time (NRT) data for the period covering the Nord Stream leaks, is viewable through the public visualisation tool (<http://rsg.rl.ac.uk/vistool>, last access 18/07/2023). NRT data is available through contacting the authors. The TOMCAT model output for this period will be made available on the Centre for Environmental Data Analysis

(CEDA) long-term data archive upon publication of this work. The ICOS methane concentrations were downloaded from ICOS Carbon portal (<https://data.icos-cp.eu/portal/>, last access 18/07/2023).

#### **Author contribution**

CW, BJK, and JJR conceptualised the study. BJK, RS, and LJV produced the satellite data. CJW, DPM, ED, WF and MPC carried out data analysis and modelling. All co-authors contributed to the design of the study and to writing the manuscript.

#### **Competing interests**

The authors declare that they have no conflict of interest.

#### **Acknowledgements**

This work was funded by the Natural Environment Research Council through its grants to the UK National Centre for Earth Observation (NCEO; NERC grant numbers NE/R016518/1 and NE/N018079/1). The IASI retrievals were produced using JASMIN, the UK collaborative data analysis facility, at the Rutherford Appleton Laboratory. The TOMCAT model simulations were carried out using ARC4, part of the High-Performance Computing facilities at the University of Leeds, UK. EUMETSAT provided data for MetOp-B IASI, MHS & AMSU-A data and ECMWF provided meteorological data used in NRT processing system and TOMCAT simulations. We thank ICOS PIs for providing their methane concentration data.

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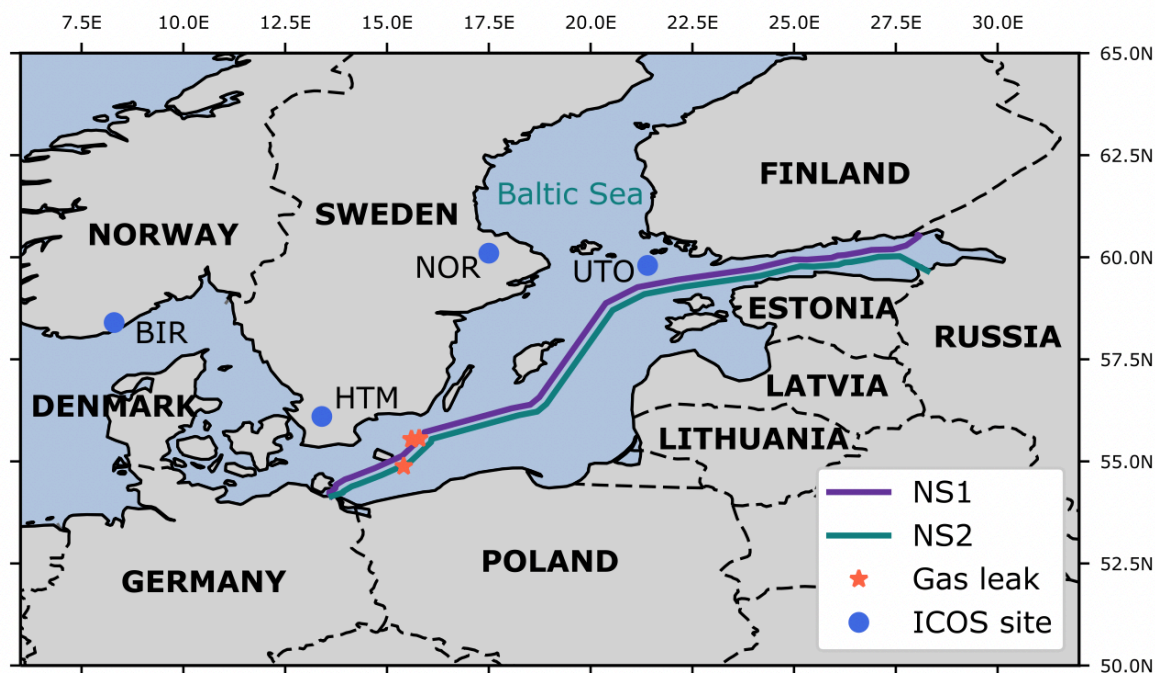


Figure 1: Map showing Nord Stream pipeline routes (teal and purple lines), gas leak locations (red stars) and *in situ* ICOS monitoring site locations (blue circles).

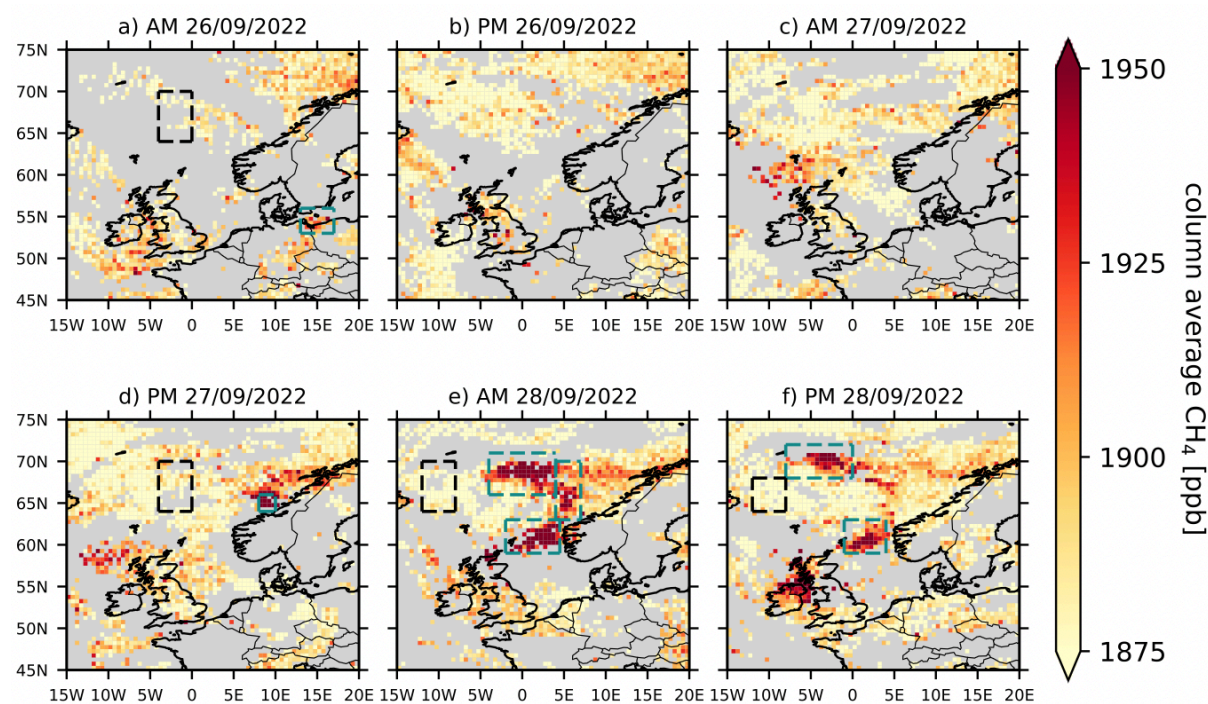


Figure 2: IASI column average CH<sub>4</sub> (ppb) for 26th - 28th September 2022. Retrievals are averaged onto  $0.25^\circ \times 0.25^\circ$  grid boxes, weighted inversely to their uncertainties for the morning and evening overpasses of each day. Black dashed boxes show 'background' regions used in the integrated mass enhancement (IME) method, whilst turquoise dashed boxes show 'enhancement' regions. Grey regions are obscured by cloud.

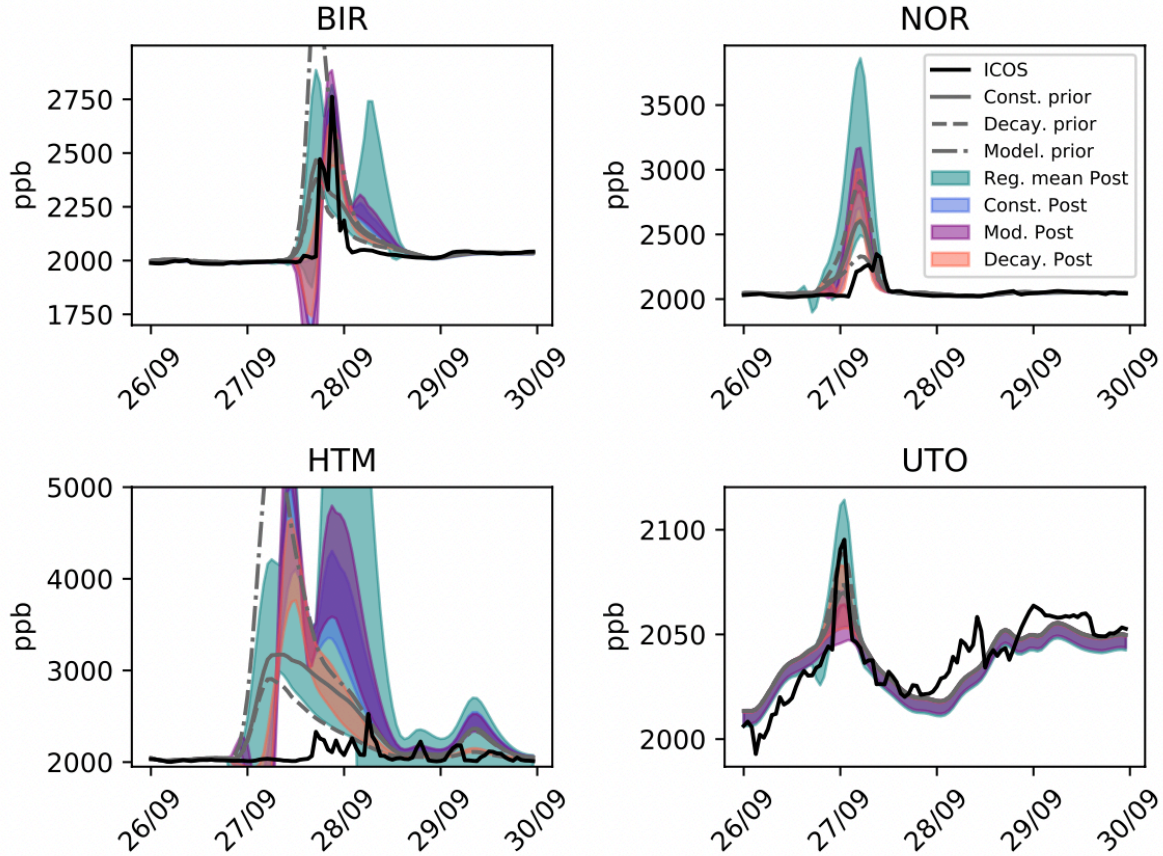


Figure 3: Observed (black line) and simulated (grey lines/coloured shading) CH<sub>4</sub> mixing ratios (ppb) at Integrated Carbon Observation System (ICOS) sites during 26<sup>th</sup> – 29<sup>th</sup> September 2022. Observations and model output are both averaged into hourly means. ICOS sites are at Birkenes, Norway (BIR), Norunda, Sweden (NOR), Hyltemossa, Sweden (HTM) and Utö, Finland (UTO). See main text and Figure 1 for further details. Grey lines show TOMCAT-simulated CH<sub>4</sub> using the three prior emission estimates, and shaded regions show the simulated min/max range for the inversions with constant prior (blue), decaying prior (red) and modelled prior (purple) optimised against individual retrievals, and for inversions optimised against the regional mean (teal). Inlet heights are the highest available at each site: 75m at BIR; 100m at NOR; 150m at HTM and 57m at UTO. Note the different y-axis ranges in each panel.



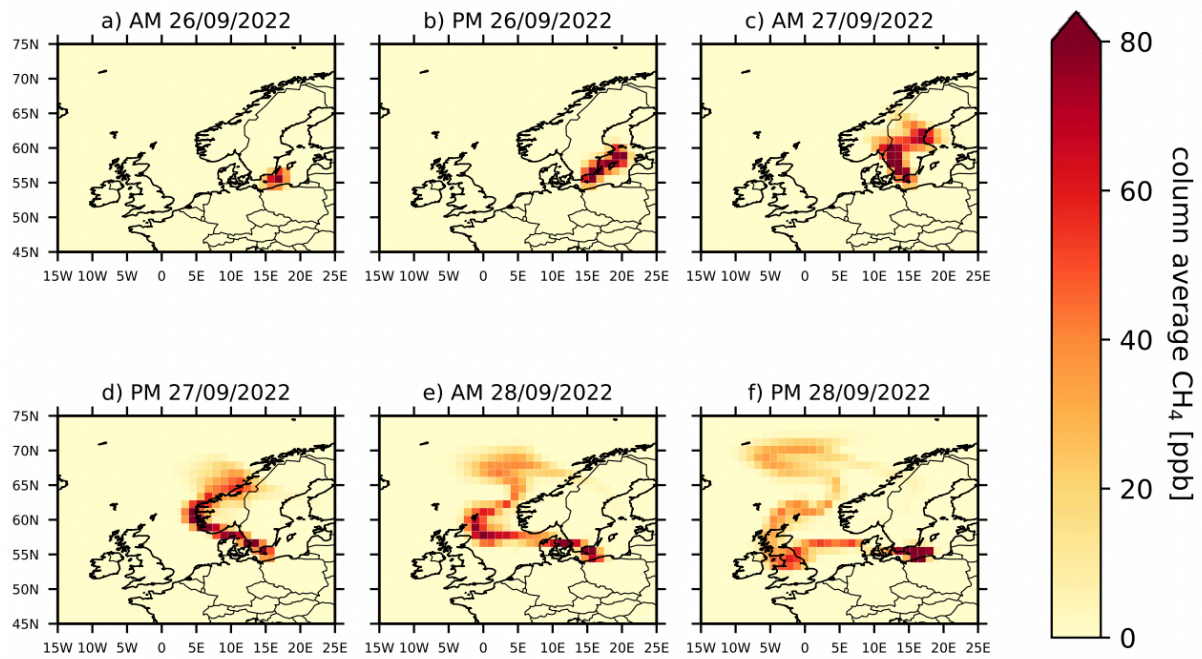


Figure 4: Simulated TOMCAT column average CH<sub>4</sub> (ppb) from Nord Stream gas leaks for 26<sup>th</sup> – 28<sup>th</sup> September 2022. Background CH<sub>4</sub> and emissions from sources other than Nord Stream are not included. Output times are matched to IASI local overpass times, but IASI averaging kernels have not been applied. Column averages are displayed on the model grid with horizontal resolution  $1.125^\circ \times 1.125^\circ$ . Emission rates from the leaks is constant at  $4.17 \text{ Gg hr}^{-1}$ , summing to 300 Gg in total over the three days.

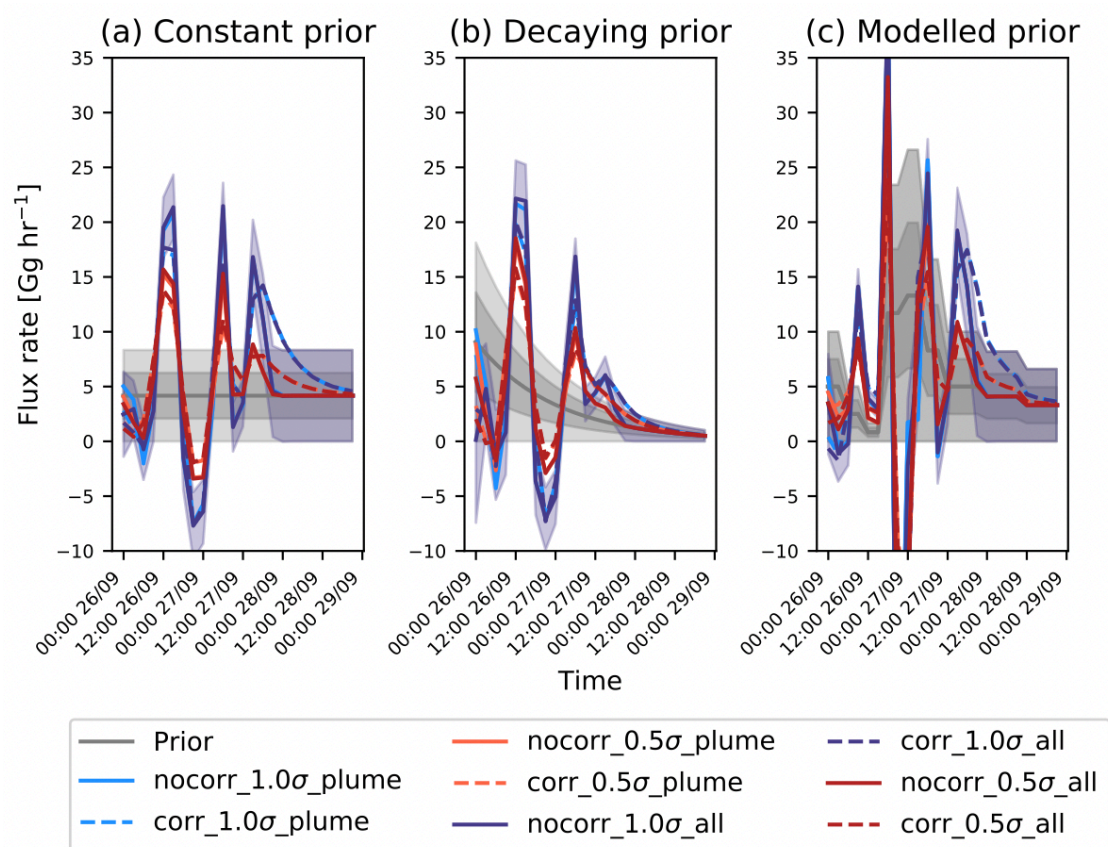
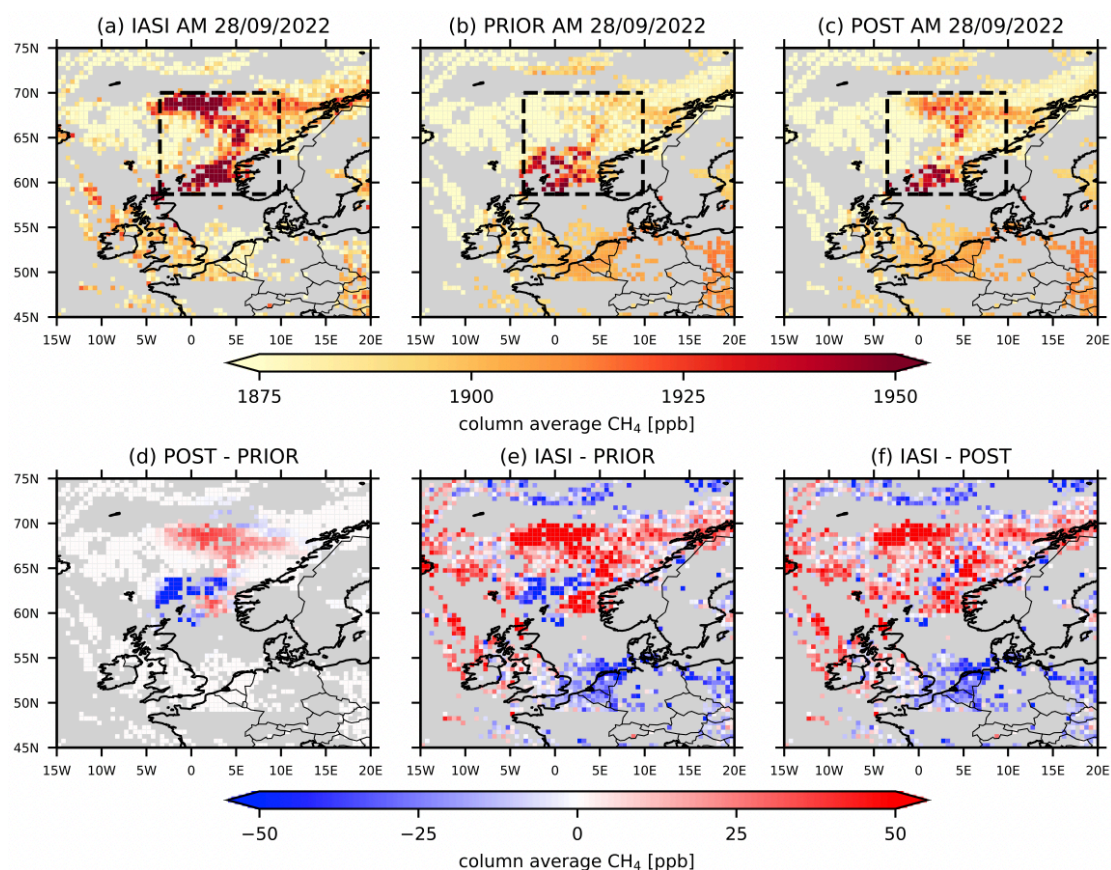


Figure 5: Prior and posterior CH<sub>4</sub> flux rates (Gg hr<sup>-1</sup>) over the first three days (September 26th - 28th) of the Nord Stream leaks based on IASI data from the morning of 28<sup>th</sup> September 2022. Prior flux rate is shown in grey, with dark grey shaded region showing the 50% prior uncertainty and the light grey shaded region showing the 100% prior uncertainty. Dashed lines show posterior inversions with prior temporal correlations imposed; solid lines show those without prior correlations. Blue lines show inversions with 100% prior uncertainty imposed; red lines show those with 50% prior uncertainty. Darker shades show inversions based on all available IASI data; lighter shades show inversions based only on IASI data from near the plume, in the region highlighted in Figure 6. Shaded blue region shows the posterior uncertainty for the 'nocorr\_1.0σ\_all' case.





**Figure 6: Column average CH<sub>4</sub> (ppb) on the morning of 28<sup>th</sup> September over the region of the Nord Stream gas leaks from (a) IASI; (b) TOMCAT using the constant prior emissions; and (c) TOMCAT using the nocorr\_1.0\_plume posterior emissions based on that prior. Also shown is the difference between the model posterior and prior (d); the difference between IASI and the model prior (e); and the difference between IASI and the model posterior (f). Retrievals and model output are averaged onto  $0.25^\circ \times 0.25^\circ$  grid boxes, weighted inversely to the observations' uncertainties. IASI averaging kernels are applied to the TOMCAT output. Black dashed line shows the 'plume' region defined in the text, used for optimising only the regional mean XCH<sub>4</sub> value.**

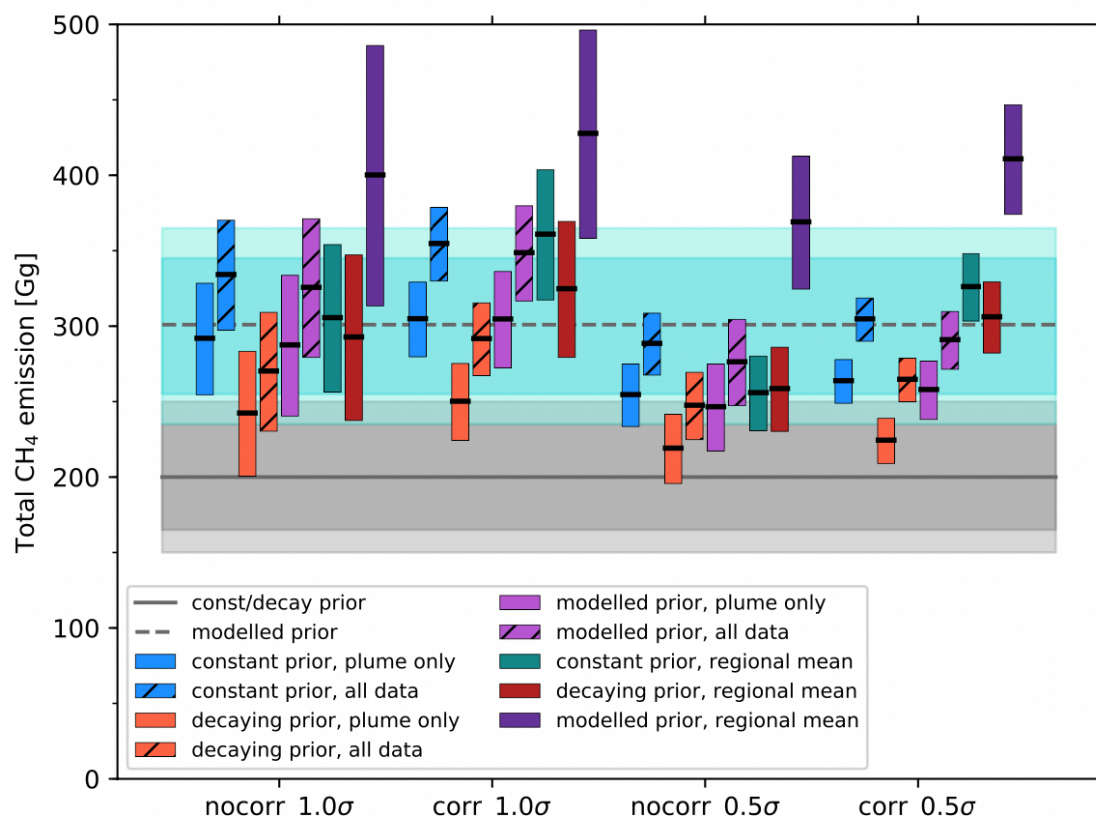


Figure 7: Total (two-day) posterior CH<sub>4</sub> emissions (Gg) from the Nord Stream leaks during 26<sup>th</sup> – 27<sup>th</sup> September based on multiple different IASI-based inverse modelling calculations. Boxes represent the 1 $\sigma$  uncertainty on each inversion, with each thick horizontal black line representing the mean. Blue boxes represent inversions with the constant prior where the model is optimised against individual IASI retrievals, whilst orange boxes are the same but for the decaying prior and pink are for the modelled prior. Teal boxes represent inversions with the constant prior where the model is optimised against the mean XCH<sub>4</sub> in the plume region, whereas red boxes are the same but for the decaying prior and purple for the modelled prior. Hatched boxes show inversions in which all IASI data is included, and unhatched boxes show inversions in which retrievals only within the plume region are included. ‘Corr’ and ‘nocorr’ refers to inversions with and without prior temporal correlations included, whilst 1 $\sigma$  and 0.5 $\sigma$  refer to inversions with 100% and 50% prior uncertainty. The grey solid line shows the prior emission total for the constant and decaying priors, with 50% and 100% 3-hour prior uncertainty shaded in dark and light grey, respectively. The grey dashed line shows the modelled prior emission, with 50% and 100% 3-hour prior uncertainty shaded in dark and light turquoise, respectively.